THE MATHEMATICAL MODELLING OF RADIONUCLIDE TRANSPORT IN THE SUBSURFACE ENVIRONMENT AROUND THE CHORNOBYL NUCLEAR POWER PLANT

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The Shelter constructed above the destroyed Unit-4 of the Chornobyl Nuclear Power Plant contains 20 MCi of nuclear fuel. More than 1000 m³ of intermediate-level radioactive water is disposed in its basement. The purpose of this work is to simulate migration of the radionuclides 90 Sr and 137 Cs from the Shelter into the subsurface environment to evaluate their migration rate and migration paths. A mathematical model accounting for the coupled transport of water and radionuclides in variably saturated media is used. The lack of suitable experimental and field studies excluded the possibility of complete validating the model. Results of the simulations will be useful for future studies of the environmental impact of the Shelter.

1. Introduction

A Shelter containment structure constructed upon the ruins of the destroyed reactor building of the Chornobyl Unit-4 was completed in November 1986. This containment was built in very high radioactive fields in the short time span of six months by using unique and remote construction methods. However, these remote methods for pouring fresh concrete did not always permit good observation and view of work, and in a number of cases high radioactive fields made welding key containment components impossible. As a result of this, openings and breaks in the walls and the roof of the Shelter are estimated to total about $1200\,\mathrm{m}^2$.

By current estimates, the Shelter contains from 135 to 180 tons of nuclear fuel with an activity of up to 20000000 Ci [1]. Within the Shelter the largest part of the nuclear fuel from the completely destroyed reactor core is contained in fuel-containing materials (FCM), and the rest – in fuel assemblies and their fragments. The FCM can be represented as lava-like materials with a glass silicate matrix containing many inclusions of various uranium compounds, finely dispersed high-level active fuel dust, and aerosols. Experimental observations show a gradual increase in the migration rate of radionuclides from the FCM caused by their permanent interaction with the water and atmosphere inside the Shelter. Contamination of the environment by radioactive substances from the Shelter may occur in two principal ways: (1) release of fuel dust into the atmosphere through the openings in the Shelter if there is a collapse of internal structures within the Shelter, and (2) migration of radionuclides through the subsurface together with the water which is inside the Shelter.

According to previous research [2,3] the Shelter contains more than 3000 m³ of intermediate-level and low-level radioactive water. More than 1000 m³ of intermediate-level radioactive water is disposed in the basement of Unit-4. Most of the contaminated water is located in the turbo-generator hall and rooms on the first level in the northern part of the Shelter [4] (see figure 1). It is also reasonable to consider such major water sources within

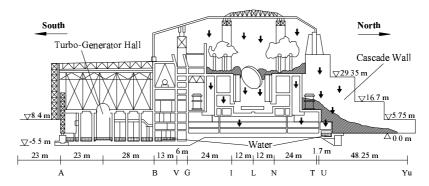


Figure 1. Schematic representation of the Chornobyl Unit-4 Shelter, water pathways and locations within the Shelter. North-south cross-section along Axis 47.

the Shelter as: precipitation, water used for dust suppression, moisture condensation onto building structures, and water leakage from the service lines of the 3rd unit of the Chornobyl NPP. Most precipitation onto the roof penetrates through cracks to the interior of the Shelter. The volume of this precipitation is approximately $1000\,\mathrm{m}^3$ per year. For dust suppression, $15\,\mathrm{m}^3$ of water are sprayed twice a month within the Shelter. Water enters the Shelter through the above mentioned openings and flows downwards into the lower regions of the reactor building where most of the fuel is located. Upon movement through the FCM and interaction with them, the water is contaminated by radionuclides. The contaminated water penetrates into the subsurface environment. The total activity of the intermediate-level radioactive water lies within the range of 2×10^{12} to $9\times10^{13}\,\mathrm{Bq/l}$. Cesium isotopes constitute the main portion of the water activity; the cesium activity varies from 1.6×10^5 to $5.5\times10^7\,\mathrm{Bq/l}$. The total $^{90}\mathrm{Sr}$ activity in the water was determined to be 3.6×10^3 to $1.1\times10^6\,\mathrm{Bq/l}$.

The Chornobyl NPP is located on the right bank of the Prypiat river. The NPP site is a flat river bench. The aquifer system is composed of fine, medium and sometimes coarse alluvial sands with interlayers of sand loam and loam. Medium sands prevail. According to field studies the hydraulic conductivity of the alluvial sands ranges from 6.0 to 15.0 m day⁻¹. These alluvial sands the thickness of which varies from 25 to 40 m are underlain by clay marls of the Kyiv Eocene suite. The marls have very low permeability and act as an aquitard. The infiltration is estimated between 80 mm and 200 mm per year [5].

The purpose of this work is to simulate migration of 90 Sr and 137 Cs through geologic media due to different scenarios of water leakage from the Shelter.

2. Model description

2.1. Water flow equations

Mathematical description of water transport in a variably saturated subsurface implies assumptions which are commonly done for this type of problems. It is assumed that fluid transport through the soil occurs in response to pressure gradients and gravitational body forces following Darcy's flow equations. The porous medium is considered to be rigid. Pressure gradients in the gas phase are so negligible that gas pressure remains effectively constant at the atmospheric pressure.

Neglecting hysteresis and temperature gradients, the movement of moisture through a saturated-unsaturated porous medium is described by

$$\frac{\partial}{\partial t}(\phi S \rho) = \frac{\partial}{\partial x_i} \left[\frac{\rho k_r k_{ij}}{\mu} \left(\frac{\partial p}{\partial x_j} + \rho g \frac{\partial z}{\partial x_j} \right) \right] + f, \tag{2.1}$$

where t is the time (T); x_i are the cartesian spatial coordinates (L); ϕ is the porosity of the medium; ρ is the fluid density (ML^{-3}) ; k_{ij} are components of the intrinsic permeability tensor (L^2) ; k_r is the relative permeability $(0 \leq k_r \leq 1)$; μ is the fluid dynamic viscosity $(ML^{-1}T^{-1})$; p is the fluid pressure $(ML^{-1}T^{-2})$; g is the gravitational constant (LT^{-2}) ; z is the elevation measured from a reference datum (L); f represents sources or sinks of fluid in the system $(ML^{-3}T^{-1})$; and S is the volumetric saturation.

Initial conditions for equation (2.1) consist of initial values of pressure, specified by

$$p(x,0) = \varphi(x),$$

where x is a spatial coordinate vector and φ is a prescribed function. Boundary conditions may be Dirichlet (type 1):

$$p(x,t) = g_1(x,t)$$
 on Γ_1 ,

or Neumann (type 2):

$$\left[k_r \frac{k_{ij}}{\mu} \left(\frac{\partial p}{\partial x_i} + \rho g \frac{\partial z}{\partial x_i}\right)\right] n_i = g_2(x, t) \quad \text{on} \quad \Gamma_2,$$

where g_1 is the prescribed pressure on the boundary Γ_1 ; n_i is the ith component of the unit vector normal to the boundary Γ_2 , directed outward; and g_2 is the prescribed outward fluid flux normal to Γ_2 . Γ_1 and Γ_2 comprise the entire boundary of the domain. Usually Γ_1 is the part of boundary at which surface-water bodies are intercepted by aquifers; Γ_2 are the boundaries across which fluid passes at a specific rate (impervious boundaries, areas of infiltration or evapotranspiration, points of injection and withdrawal, etc.).

2.2. Transport equations

Let c_l represent the volumetric radionuclide activity in the aqueous phase (ML^{-3}) ; c_s – the volumetric exchangeable sorbet activity in the solid phase (ML^{-3}) ; and c_f – the volumetric radionuclide activity which is fixed in the mineral lattice (ML^{-3}) . Assuming that the ion-exchange reaction transferring activity between aqueous and solid exchangeable phases has a sufficiently short timescale, we will consider this reaction as instantaneous which

is described by a linear equilibrium isotherm. Transfers of activity to and from the fixed phase are modelled by first-order rate constants α_{sf} and α_{fs} , respectively, as shown in figure 2 [6].

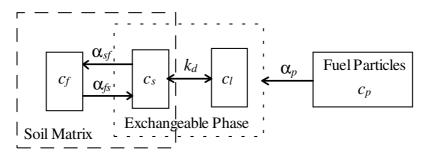


Figure 2. Schematic representation of the kinetic sorption model.

The leaching of radionuclides from the fuel particles is modelled by a first-order equation:

$$\frac{\partial c_p}{\partial t} = -(\alpha_p + \lambda)c_p,$$

$$c_p(x, 0) = c_p^0(x),$$

where c_p is the radionuclide activity in the fuel particles per soil solid volume (ML^{-3}) ; α_p is the first-order constant of radionuclide leaching from fuel particles (T^{-1}) ; λ is the radionuclide decay constant (T^{-1}) ; c_p^0 is the initial radionuclide activity in the fuel particles (ML^{-3}) .

Ignoring chemical diffusion in the solid phase, an equation for adjectivedispersive transport of radionuclide activity in the exchangeable phase in variably saturated media may be written in the form:

$$\frac{\partial c}{\partial t} = \frac{\partial}{\partial x_i} \left(\phi S D_{ij} \frac{\partial c_l}{\partial x_j} \right) - \frac{\partial}{\partial x_i} (v_i c_l) - \lambda c + (1 - \phi_T) (\alpha_p c_p + \alpha_{fs} c_f - \alpha_{sf} c_s), \tag{2.2}$$

where c is the radionuclide activity per total volume (ML^{-3}) . D_{ij} is the dispersion tensor defined by

$$\phi SD_{ij} = \alpha_T \mid v \mid \delta_{ij} + (\alpha_L - \alpha_T) \frac{v_i v_j}{\mid v \mid} + \phi S \tau D_0 \delta_{ij},$$

where α_L and α_T are the longitudinal and transverse dispersiveness (L), respectively; δ_{ij} is the Kronecker delta; |v| is the magnitude of the Darcy velocity; D_0 is the molecular diffusion coefficient (L^2T^{-1}) ; and τ is the tortuosity.

The total radionuclide activity is defined by:

$$c = \phi S c_l + (1 - \phi_T) c_s,$$

where ϕ_T is the total porosity of the medium; and volumetric phase activities are interrelated through the solid-aqueous partition coefficient k_d , according to the relationship

$$k_d = \frac{c_s \rho}{c_l \rho_s},$$

where ρ_s is the soil solid density (ML^{-3}) .

Initial conditions for equation (2.2) consist of the specification of initial radionuclide activity in the aqueous phase:

$$c_l(x,0) = c_0(x).$$

Boundary conditions for inflow boundaries are Dirichlet (type 1):

$$c_l = c_1$$
 on Γ_3 ,

or Cauchy (type 3):

$$\left(-\phi S D_{ij} \frac{\partial c_l}{\partial x_j}\right) n_i = v_b c_b \quad \text{on} \quad \Gamma_4,$$

where c_1 is the prescribed radionuclide activity on the boundaries Γ_3 ; n_i is the unit vector normal to boundaries, directed inward; v_b is the prescribed fluid flux normal to the boundaries Γ_4 ; and c_b is the prescribed radionuclide activity associated with the prescribed fluid flux.

activity associated with the prescribed fluid flux.

At outflow boundaries the assumption is often made that transport across the boundaries, according to [7], occurs by advection only and the Neumann (type 2) boundary condition is specified by

$$\frac{\partial c_l}{\partial x_i} = 0 \quad \text{on} \quad \Gamma_5,$$

where $\Gamma_3 + \Gamma_4 + \Gamma_5 = \Gamma$ is the boundary of the domain. The transport of activity in the fixed phase is described by:

$$\frac{\partial c_f}{\partial t} = \alpha_{sf} c_s - \alpha_{fs} c_f - \lambda c_f,$$
$$c_f(x, 0) = c_f^0(x),$$

where c_f^0 is the initial activity in the fixed phase.

3. Simulation results

Numerical simulations of the ⁹⁰Sr and ¹³⁷Cs migration from the Shelter into and through the subsurface environment were conducted for a two-dimensional vertical cross-section selected along the streamlines of the groundwater flow. These streamlines were obtained from field observations of groundwater level distribution in the vicinity of the Chornobyl NPP [5]. The geology of this cross-section is shown in figure 3.

The MSTS code [8,9] adapted to implement the model described above was used for the numerical simulations. The flow domain was discretized by a non-uniform rectangle mesh, the nodes of which were spaced at 1 m - 100 m intervals in the horizontal direction, and from 0.2 to 3 m in the vertical direction. Time steps of 30 days were used. Values for the hydraulic and species properties used in the simulations are given in table 1. Values of the first-rate constants $\alpha_{sf}, \ \alpha_{fs}$ were taken similar to the experimentally obtained values of these parameters for the upper soil layer of the left bank floodplain of the Prypiat river. Parameters α_{sf} and α_{fs} vanished for the

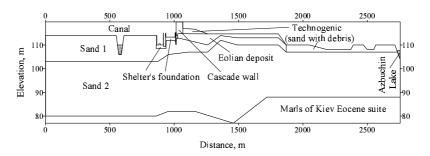


Figure 3. Schematic representation of the geologic section.

Table 1. Values of parameters used in numerical simulations.

Soil or Rock	$K \ { m m day^{-1}}$	$lpha_L$ m	$lpha_T$ m	k_d^{Sr} l kg $^{-1}$	k_d^{Cs} l kg ⁻¹	$lpha_{sf}^{Cs} \ \mathrm{day}^{-1}$	$\begin{array}{c} \alpha_{fs}^{Cs} \\ \mathrm{day}^{-1} \end{array}$
Sand 1 Sand 2 Eolian deposit Technogenic Foundation Cascade wall	9.0 10.0 13.0 11.0 0.001 1.0	1.2 1.2 1.2 1.2 1.0 1.0	$0.1 \\ 0.1 \\ 0.1 \\ 0.1 \\ 0.1 \\ 0.1 \\ 0.1$	1.0 1.0 1.0 1.0 0.5 0.5	5.2 5.2 5.2 5.2 5.2 5.2	0.001 0.001 0.001 0.001 0.001 0.001	0.00005 0.00005 0.00005 0.00005 0.00005

case of 90 Sr migration in the subsurface. The coefficient of radionuclide leaching from fuel particles was 3.2×10^{-4} day $^{-1}$ [1].

The prescribed values of the hydraulic head were specified along the left and right boundaries, the top of the Shelter's foundation, and the boundaries on which surface water bodies are intercepted by the aquifer. At the ground surface the liquid flow rate was equal to the infiltration rate or evapotranspiration rate. The annual amount of infiltration was 80 mm. The marls were considered as impervious. The water depth was assumed to be 2 m in the turbo-generator hall, and 1.6 m in rooms on the first level in the northern part of the Shelter. Seasonal fluctuations of the water level in the canal varied between 109.5 m and 110.2 m; in Azbuchin Lake – between 106.0 m and 106.8 m; and for the hydraulic head at the left boundary – between 109.5 m and 110.0 m [5]. These seasonal fluctuations of the water level in surface water bodies and the variability over time of precipitation and evapotranspiration were taken into account in the simulations.

Before the Chornobyl accident a special drainage system operated to prevent underflooding of the Chornobyl NPP building's foundation. After the Chornobyl accident this drainage system was damaged and shut down. As a result of that a groundwater level increase was observed until 1994. The nowadays groundwater level below the Shelter varies from 109.5 m to 110.5 m depending on seasonal groundwater fluctuations and precipitation. The groundwater increase after the Chornobyl accident was not taken into account in the simulations. The initial condition for the liquid phase is illustrated in figure 4.

A zero radionuclide flux condition was employed at all the boundaries,

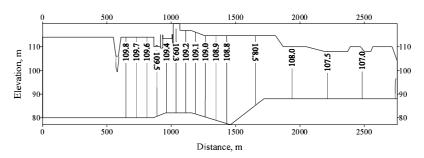


Figure 4. Initial hydraulic head distribution.

excluding the Shelter basement and the boundaries that contact surface water bodies. The 90 Sr activity of water from the basement was 2.3×10^6 Bq/l, and 2800 Bq/l in Azbuchin Lake [5]. Similarly, 137 Cs activity in the basement water and Azbuchin Lake was 7.1×10^7 and 84 Bq/l, respectively. From 1986 to 1995, the radionuclide activity of canal water was determined from the annual average activity of water samples from the Cooling pond. The radionuclide activity of canal water after 1995 was specified as the radionuclide activity of canal water in 1995. For the area surrounding the Chornobyl NPP, land surface contamination by 137 Cs in 1995 was from 3.8×10^{-3} to 1.9×10^{-4} Ci per kg of soil [2]. 90 Sr land surface contamination was equal to one half of 137 Cs and surface contamination. Just after the accident, exchangeable 90 Sr was 5% of the total 90 Sr activity and exchangeable 137 Cs was 10% of the total 137 Cs activity. The subsurface was assumed to be uncontaminated before the accident. The date of the Chornobyl accident was taken as the starting date for the simulations.

There are no reliable data on water leakage beyond the bounds of the Shelter. However, analysis of water pathways and water locations within the Shelter shows that there is a possibility of water leakage from the turbogenerator hall, northern rooms of the Shelter, and through the cascade wall. The cascade wall was filled with concrete, boxed radioactive materials collected from the roof and surroundings, building debris, and structural metals. So, it has a lot of openings and breaks through which part of precipitation onto the cascade wall penetrates into the subsurface. Therefore, in order to estimate migration rates of radionuclides from the Shelter through the subsurface, numerical simulations were conducted for the following cases:

- Case 1. Water leakage occurs from rooms on the first level in the northern part of the Shelter.
- Case 2. In addition to case 1, one half of the precipitation onto the cascade wall penetrates into the outside of the Shelter.
- Case 3. Water from the turbo-generator hall, rooms in the northern part of the Shelter, and a half of the precipitation onto the cascade wall penetrates into the subsurface environment.

The effects of various water leakage scenarios on radionuclide migration from the Shelter through the subsurface are compared with case 2. Figures 5 and 6 show the predicted activity distribution of 90 Sr and 137 Cs radionuclides in the aqueous phase in 1995 and 2045 for case 2.

The influence of different water leakage into the outside of the Shelter on migration of ⁹⁰Sr through the subsurface is shown in figures 7-9. Analysis of

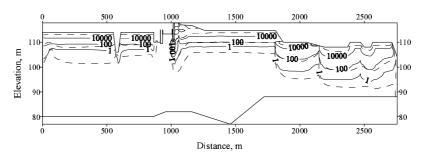


Figure 5. Activity of ⁹⁰Sr (Bq/l) in the aqueous phase in (—) 1995 and (---) 2045 for case 2.

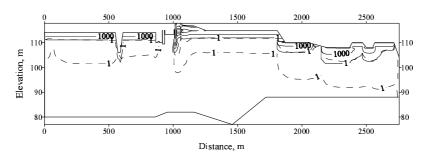


Figure 6. Activity of ¹³⁷Cs (Bq/l) in the aqueous phase in (—) 1995 and (---) 2045 for case 2.

the results indicates the increase of the groundwater level and radionuclide migration rate below the Shelter for case 3. The effect of water penetration through the cascade wall on radionuclide migration below the Shelter is negligible.

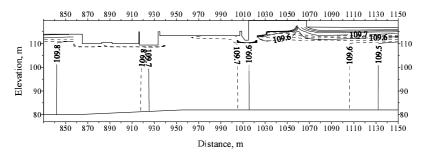


Figure 7. Summer distribution of hydraulic head below the Shelter for (—) case 2 and (---) case 3.

Comparisons of predicted activity-depth profiles of 90 Sr and 137 Cs radionuclides in the aqueous phase with the measurements in 1995 are given in figures 10 and 11. These predicted activity-depth profiles were selected at point x=1075 m which approximately corresponds to the location of wells 1-g, 2-g and 4-g in the northern part of the Chornobyl NPP building area. The annual mean values of radionuclide activity obtained on measurements in wells 1-g, 2-g, 4-g and 6-g [2,3] are denoted by circles. A triangle denotes the radionuclide activity detected in well C10 in 1995 [5]. Well C10

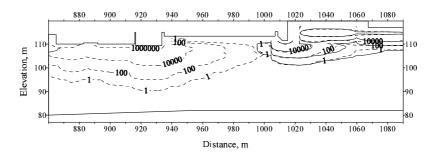


Figure 8. Activity of ⁹⁰Sr (Bq/l) in the aqueous phase below the Shelter in 1995 for (—) case 2 and (---) case 3.

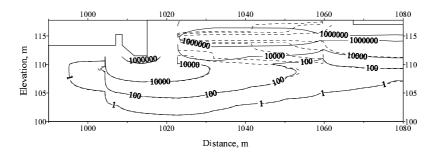


Figure 9. Activity of 90 Sr (Bq/l) in the aqueous phase below the Shelter in 1995 for (—) case 2 and (---) case 1.

Aqueous activity (Bq/l) of Sr-90

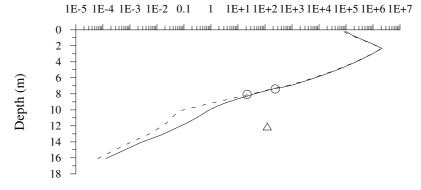


Figure 10. Comparison of predicted and measured activity-depth profiles of $^{90}\mathrm{Sr}$ in the aqueous phase in 1995. Model-generated profiles were obtained for (—) case 2 and (---) case 3. \triangle and \bigcirc denote mean values of radionuclide activity detected in well C10 and in wells 1-g, 2-g, 4-g and 6-g, respectively.

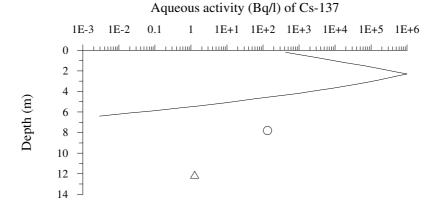


Figure 11. Comparison of predicted and measured activity-depth profiles of $^{137}\mathrm{Cs}$ in the aqueous phase in 1995. Model-generated profile was obtained for (—) case 2. \triangle and \bigcirc denote mean values of radionuclide activity detected in well C10 and in wells 1-g, 2-g, 4-g and 6-g, respectively.

is located approximately 150 m east of wells 1-g and 4-g.

4. Conclusions

This work has concentrated on modelling ⁹⁰Sr and ¹³⁷Cs migration from the Shelter into and through the subsurface environment. A two-dimensional mathematical model accounting for the coupled transport of water and radionuclides in variably saturated porous media was used. The hydraulic conductivity, the longitudinal and transverse dispersiveness are sensitive parameters for radionuclide transport modelling. The key parameters for radionuclide exchange between soil water and soil matrix are the distribution coefficient or k_d , and first-order rate constants α_{sf} and α_{fs} . Due to the limited availability of field data, these parameters were assigned by experts. As a result of that only a qualitative radionuclide transport modelling was achieved. However, modelling results presented in this work illustrate the basic effects of different possible water leakage into the outside of the Shelter on the migration rate of radionuclides through the subsurface. The simulation results showed that, so far, the groundwater contamination detected in the observation wells is only defined by the surface land contamination. Much more quantitative testing must be conducted to evaluate the environmental impact of the Shelter on the subsurface contamination.

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МАТЕМАТИЧНЕ МОДЕЛЮВАННЯ ПЕРЕНОСУ РАДІОНУКЛІДІВ У ПРИПОВЕРХНЕВОМУ СЕРЕДОВИЩІ НАВКОЛО ЧОРНОБІЛЬСЬКОЇ АЕС

С.Л.Ківва

Саркофаг, що споруджено над зруйнованим четвертим блоком Чорнобильської АЕС, містить 20 МКи ядерного палива. Більш ніж 1000 м³ середньоактивної води знаходиться на його нижніх поверхах. Метою даної роботи було моделювання міграції радіонуклідів 90 Sr та ¹³⁷ Cs з Саркофагу в геологічне середовище для оцінки їх швидкості та шляхів міграції. Для цього використовувалась сумісна модель руху вологи та радіонуклідів в насичено-ненасиченому пористому середовищі. Недостатня кількість польових і лабораторних досліджень не дозволила повністю верифікувати модель. Результати моделювання можуть бути використані в подальших дослідженнях з оцінки впливу Саркофагу на забруднення оточуючого середовища.